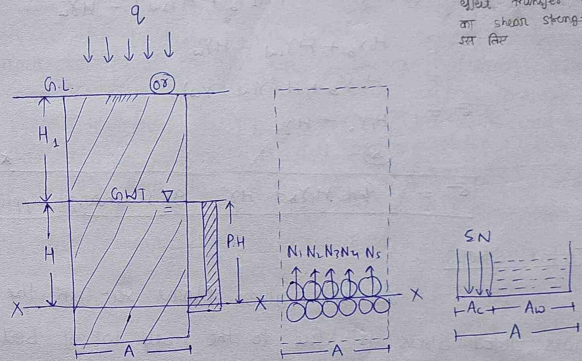


Lecture 5
7/11/19

CHAPTER-3

load लगाने से grain to grain
जुड़ कर transfer होगा इसके बाद
का shear strength to हाता है
सा कि



$$\begin{aligned} \sigma &= \frac{\text{total wt}}{\text{Area}} = \frac{\text{wt of (solids + water)}}{\text{Area}} \\ \sigma &= \frac{V \gamma'}{\text{Area}} = \frac{\text{depth} \times \text{Area} \times \gamma'}{\text{Area}} = \text{depth} \times \gamma' \\ \sigma &= q + \text{depth} \times \gamma' \end{aligned}$$

On Plane (X-X) $\sigma = H \gamma_b + H \gamma_{sat}$

$\sigma = q + H \gamma_b + H \gamma_{sat}$ (if surcharge is applied)

total stress (σ) = Eff stress + $H \gamma_{sat}$

$$\sigma = \bar{\sigma} + u$$

(u) Pore water pressure = $\frac{\text{wt of water column}}{\text{Area}}$

$$u = \frac{\text{vol} \times \gamma_w}{\text{Area}} = \frac{A_w \cdot H \cdot \gamma_w}{A_w} = H \gamma_w$$

$$u = \text{Pressure Head} \times \gamma_w = H \gamma_w$$

Effective stress = total stress - Pore water pressure

$$\bar{\sigma} = (\sigma - u) = (H_b \gamma_b + H \gamma_{sat}) - H \gamma_w$$

$$= H_b \gamma_b + H \gamma_{sat} - H \gamma_w$$

$$= H_b \gamma_b + H (\gamma_{sat} - \gamma_w)$$

$$\bar{\sigma} = H_b \gamma_b + H \gamma'$$

②

$$\bar{\sigma} = q + H_b \gamma_b + H \gamma' \quad \text{if surcharge applied}$$

* Total Stress

At any given plane sec in the

soil mass total stress is due to the self wt of soil (wt of (solids + water + air)) all due to the applied overburden pressure (uniform surcharge)

→ This total stress further consist of two different component

- 1) Effective stress (intergranular pressure)
- 2) Pore water pressure (Neutral pressure)

Effective stress ($\bar{\sigma}$)

It is the stress which is being transferred in the soil mass by grain to grain contact which tend to force the particles to come into closure state of contact resulting in its decreased void ratio, increase degree of denseness and mobilization of shear strength.

These stresses are also termed as inter-granular stresses.

Pore water pressure (u)

It is the pressure which is being transmitted by pore field and is equal to the wt of water column above the concern sec in soil mass. This water pressure can also be determined using Piezometer (Pressure head $\times \gamma_w$) this water pressure acts all around the soil solid Hence does not tend to force the soil solid into closure state of contact. It does not have any shear component. It is also termed as Neutral pressure

Approach II

$$\text{Total Stress } (\sigma) = \frac{\text{total force}}{\text{area}} = \frac{\Sigma N + U A_w}{A} = \frac{\Sigma N}{A} + \frac{U A_w}{A}$$

$$\therefore A = A_c + A_w$$

$$A_c \ll A_w$$

$$A \approx A_w$$

$$\sigma = \frac{\Sigma N}{A} + \frac{U A_w}{A}$$

$$\text{Total stress} = \bar{\sigma} + u$$

$$\text{effective stress} = \frac{\Sigma N}{A}$$

Note ①

$$\text{Contact stress} = \frac{\Sigma N}{A_c}$$

$$\text{Eff. stress} = \frac{\Sigma N}{A}$$

$$\therefore A_c \ll A$$

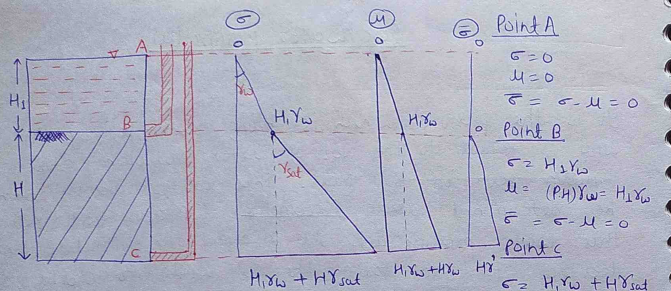
$$\frac{\Sigma N}{A_c} \gg \frac{\Sigma N}{A}$$

Note ② (consolidation/settlement)

(shear strength and bearing capacity & the funn of eff. stress)

Case-①

Submerge soil Mass (GWT is above the ground level)

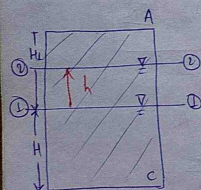


Note

⇒ If water table is above the ground level then variation of g_w water table will not affect the effective stress.

⇒ But if water table is below the ground level then variation of water table will affect the effective stress

Case-① change in effective stress due to variation of GWT level
 (GWT is below the ground level)



effective stress at c when GWT is at ①-①

$$(\bar{\sigma}_c)_{1,1} = \sigma - u = (H_1 \gamma_b + H \gamma_{sat}) - H \gamma_w$$

$$= H_1 \gamma_b + H \gamma'$$

effective stress at c when GWT is at ②-②

$$(\bar{\sigma}_c)_{2,2} = \sigma - u = [(H_1 - h) \gamma_b + (H + h) \gamma_{sat}] - (H + h) \gamma_w$$

$$= (H_1 - h) \gamma_b + (H + h) \gamma'$$

Change in effective stress

$$\Delta \bar{\sigma} = (\bar{\sigma}_c)_{2,2} - (\bar{\sigma}_c)_{1,1} = [(H_1 - h) \gamma_b + (H + h) \gamma'] - [H_1 \gamma_b + H \gamma']$$

$$\Delta \bar{\sigma} = -h \gamma_b + h \gamma'$$

$$\Delta \bar{\sigma} = h(\gamma' - \gamma_b) = -ve$$

Note If GWT is below the G.L

GWT rises → effective stress decreases

GWT lowered → effective stress increases

Note ① When water table is lowered due to pumping out of water then :- immediately after the lowering soil will be treated as saturated in lowered water table zone

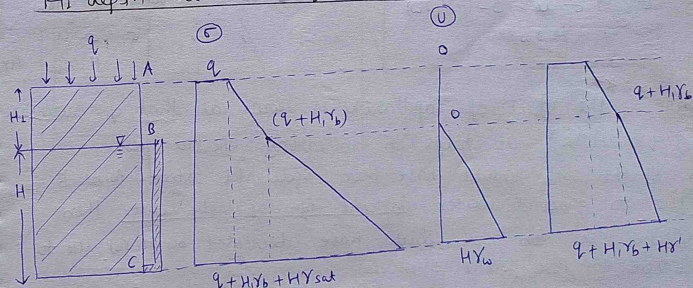
$$\Delta \bar{\sigma} = (\gamma_{sat} - \gamma') h$$

② In long term (after a long time soil will achieve the bulk cond in lower water table zone

$$\Delta \bar{\sigma} = (\gamma_{bulk} - \gamma') h$$

LOGIC

Case-③ Soil Mass with uniform surcharge (GWT is H_1 depth below the ground level :-



Point A

$$\sigma = q$$

$$u = 0$$

$$\bar{\sigma} = \sigma - u = q$$

Point B

$$\sigma = q + H_1 \gamma_b$$

$$u = 0$$

$$\bar{\sigma} = \sigma - u = q + H_1 \gamma_b$$

Point C

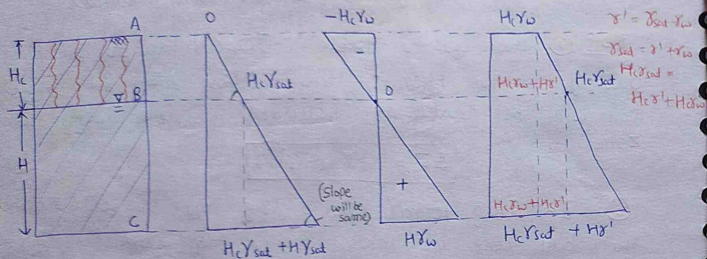
$$\sigma = q + H_1 \gamma_b + H \gamma_{sat}$$

$$u = (pH) \gamma_w = H \gamma_w$$

$$\bar{\sigma} = \sigma - u$$

$$= q + H_1 \gamma_b + H \gamma'$$

Case ④ Soil mass with capillarity fringes (soil is saturated above GWT under capillary action) gmp



Point A

$$\begin{aligned}\sigma &= 0 \\ u &= -H_c \gamma_w \\ \bar{\sigma} &= \sigma - u = 0 - (-H_c \gamma_w) \\ &= H_c \gamma_w\end{aligned}$$

Point B

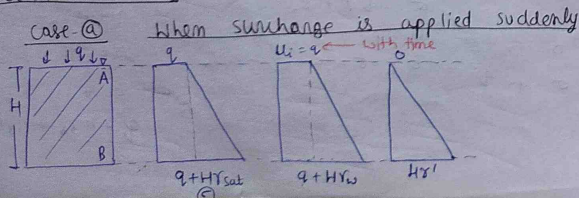
$$\begin{aligned}\sigma &= H_c \gamma_{sat} \\ u &= 0 \\ \bar{\sigma} &= \sigma - u = H_c \gamma_{sat}\end{aligned}$$

Point C

$$\begin{aligned}\sigma &= \gamma_{sat} H_c + H_c \gamma_{sat} \\ u &= H_c \gamma_w \\ \bar{\sigma} &= \sigma - u = H_c \gamma_{sat} + H_c \gamma_w\end{aligned}$$

Note *** the effect of capillarity is same as that of surcharge ($q = H_c \gamma_w$). It helps in increasing effective stress.
→ If the soil would have been saturated above point B due to ground water instant of capillary water then effective stress in soil would have decreased by ($H_c \gamma_w$) due to increased pore water pressure.

Case-⑤ Application of uniform surcharge (GWT is at ground level)



Point A

$$\begin{aligned}\sigma &= q \\ u &= 0 + u_i = 0 + q \\ \bar{\sigma} &= \sigma - u = 0\end{aligned}$$

Note

at $t = 0$

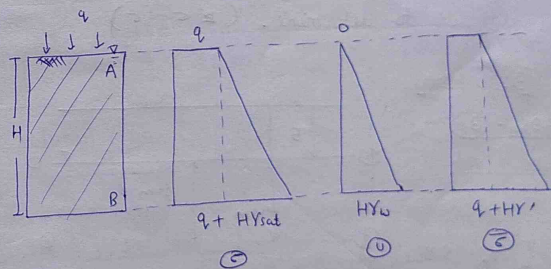
$$F = K \frac{\partial u}{\partial x} \quad (\text{at } t \text{ tends to } 0)$$

$$u_i = q$$

at any time (t)

$$\begin{aligned}\downarrow u_i \\ \uparrow F &= K \frac{\partial u}{\partial x}\end{aligned}$$

Case ⑥ When surcharge is applied gradually (GWT at ground level)



Point A

$$\begin{aligned}\sigma &= q \\ u &= 0 + u_i = 0 \\ \bar{\sigma} &= \sigma - u = q\end{aligned}$$

Point B

$$\begin{aligned}\sigma &= q + H \gamma_{sat} \\ u &= H \gamma_w + u_i = H \gamma_w \\ \bar{\sigma} &= \sigma - u = H \gamma_{sat}\end{aligned}$$

Note When W.T. present at ground level and surcharge is applied gradually it is being carried by soil solid.
→ If surcharge is applied suddenly then firstly it will be taken by water in it due to which excess pore water developed that causes water to seep out during which it transfers a part of surcharge on the solid soil solid and when excess pore water pressure is being dissipated. When total surcharge is being carried by soil solid.

Case-6 Partially saturated Soil

1) If soil is partially saturated then air is also present in the soil along with the water. Hence, in the analysis of effective stress pore air pressure should also be considered along with pore water pressure.

Bishop and Lambe

$$\bar{\sigma} = \sigma - U_a + X(U_a - U_w) \quad \left[X = \frac{A_w}{A} \right]$$

U_a = Pore air pressure

U_w = Pore water pressure

X = fraction of the area of soil occupied by the water

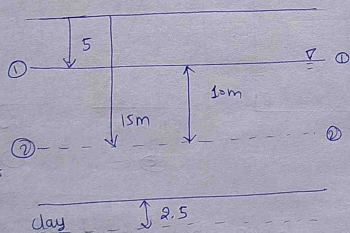
$$= \frac{[A_w]}{[A]}$$

Note In partially saturated cond pore pressure is neglected but it is difficult to determine. ($\bar{\sigma} = \sigma$)

Lecture 5 [3:30 hour] 8/11/19

Page 37 Qn

① When GWT is at ①-①



$$\begin{aligned} \bar{\sigma} &= \sigma - U = [5\gamma_s + 15\gamma_{sat}] - 17.5\gamma_w \\ &= 5 \times 18 + 15 \times 20 + 2.5 \times 20 \\ &\quad - 17.5 \times 9.81 \\ &= 268.325 \end{aligned}$$

② When GWT is at ②-②

$$\begin{aligned} \bar{\sigma} &= 15\gamma_w + 5(\gamma_{sat})_s + 2.5(\gamma_{sat})_c - 7.5\gamma_w \\ &= 15 \times 18 + 5 \times 20 + 2.5 \times 20 - 7.5 \times 9.81 = 346.425 \\ \Delta \bar{\sigma} &= 346.425 - 268.325 = 78.1 \text{ kN/m}^2 \end{aligned}$$

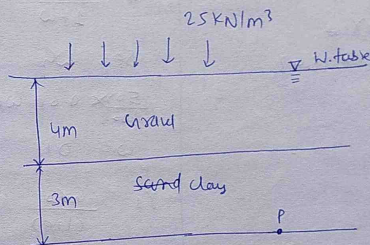
III

$$\Delta \bar{\sigma} = (\gamma_s - \gamma')h = (18 - (20 - 9.81)) \times 10 = 78.1 \text{ kN/m}^2$$

Solution (35) Page (2)

$$\gamma_{\text{clay}} = 19 \text{ kN/m}^3$$

$$\gamma_{\text{sand}} = 20 \text{ kN/m}^3$$



$$\begin{aligned} \sigma &= 162 \\ U &= 190 \\ \text{and } \bar{\sigma} &= 52 \text{ kPa} \end{aligned}$$

$$\sigma = 4 \times (25) + 3 \times (19) + 9$$

$$\sigma = 25 + 4 \times (20) + 3 \times (19)$$

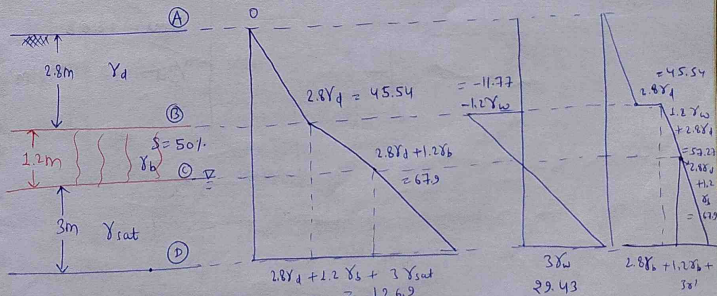
$$\sigma = 162$$

$$U = 0 + 4 \times 10 + 7 \times 10$$

$$U = 40 + 70 = 110$$

$$\begin{aligned} \bar{\sigma} &= \sigma - U \\ &= 162 - 110 = 52 \end{aligned}$$

Solution (36) Page (2)



At Point A

$$\begin{aligned}\sigma &= 0 \\ U &= 0 \\ \text{and } \bar{\sigma} &= 0\end{aligned}$$

At Point B

$$\begin{aligned}\sigma &= 2.8 \times (\gamma_d) = 2.8 \times 16.56 = 46.375 \\ U &= 2.8 \times (-1.2 \times (\gamma_w)) = -12 \\ \bar{\sigma} &= 46.375 + 12 = 58.375\end{aligned}$$

At Point C

$$\begin{aligned}\sigma &= 2.8 \times \gamma_d + 1.2 \times \gamma_{sat} \\ &= 2.8 \times 16.56 + 1.2 \times 18.43 \\ &= 68.48\end{aligned}$$

At Point D

Solution (37)

$$\gamma_d = \frac{\gamma_b}{1+w}$$

$$Se = wG$$

$$\frac{Se}{100} \times 0.6 = w \times 2.65$$

$$w = 0.11$$

$$\gamma_d = \frac{\gamma_{sw}}{(1+e)}$$

$$\gamma_d = \frac{26.5 \times 10}{(1+0.11)}$$

$$\gamma_d = 16.56 \text{ kN/m}^3$$

$$\gamma_{sat} = \frac{(G+Se) \gamma_w}{(1+e)}$$

$$\gamma_{sat} = 18.43$$

$$\gamma_{sat} = 18.43$$

$$\gamma_{sat} = 19.93 \text{ kN/m}^3$$

CHAPTER-4

PERMEABILITY

Permeability is the property of soil by virtue of which soil allows the flow of fluid through it.

⇒ It is also termed as hydraulic conductivity

⇒ Permeability of coarse grain soil is more than that of fine grain soil.

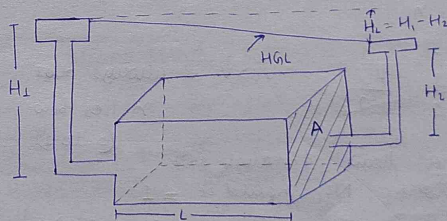
Type of soil

- 1) Gravel
- 2) Sand
- 3) Silt
- 4) clay

Permeability

- > 1 cm/sec
- $1 - 10^3$ cm/sec
- $10^{-3} - 10^7$ cm/sec
- < 10^{-7} cm/sec

DARCY'S LAW



Hydraulic gradient =

$$\frac{\text{Seepage Head}}{\text{length}} = \frac{\text{Head loss}}{\text{length}}$$

$$i = \frac{H_1 - H_2}{L} = \frac{H_L}{L}$$

As per Darcy's law for laminar flow in saturated soil mass, velocity of flow is directly proportional to hydraulic gradient.

$$V \propto i$$

$$V = K i^{**}$$

$k \rightarrow$ coefficient of permeability ($\frac{m}{sec}$, $\frac{cm}{sec}$)

Discharge = area velocity

$$Q = K i A \quad \text{m}^3/\text{sec}$$

Limitations

1) The velocity of flow considered above the avg or discharge velocity but total area of cb is considered but in actual flow takes place through the interconnectivity void of medium, area of which is much less than total area of cb of flow. Hence actual / true / seepage velocity of flow is much less than avg discharge velocity considered above.

$$\text{Discharge (Q)} = \frac{\text{Volume}}{\text{time}} = A \times V = A \times V_s$$

$$\frac{V}{V_s} = \frac{A \times V}{A \times V_s} = \frac{A \times V}{A \times \frac{V}{n}} = \frac{\text{Volm of Voids}}{\text{Total Volm}} = \text{Porosity (n)}$$

$$V = V_s \cdot n$$

Seepage velocity

$$\frac{V}{V_s} = \frac{V}{V_s} = \frac{K i}{n}$$

$V_s > V$ $\left\{ \begin{array}{l} \text{always true} \\ n \text{ lies b/w } 0 \text{ to } 1 \end{array} \right.$

Note For laminar flow through soil mass Reynold's No. should be ≤ 1

(4.2) METHODS TO DETERMINE PERMEABILITY

1) LAB METHODS

- At constant Head permeability test
- Variable Head permeability test
- Capillarity Head permeability test

(2) Field Method

- Dupit's theory
- Thiem's theory

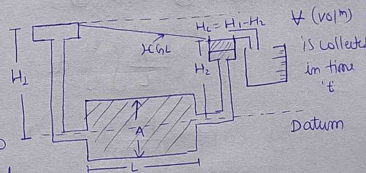
(3) Indirect Method

- Kozney - carman eqn
- Allen Hazen eqn
- consolidation eqn
- Terzaghi's eqn

A) Constant Head permeability test

* Suitable for coarse grained soil

Substantial for coarse grained soil: Substantial discharge can be collected in small time



* Water is allowed to flow under constant head and let volume V is flowing through soil mass in time t

$$\Rightarrow \text{Discharge (Q)} = \frac{\text{Volume}}{\text{time}} = \frac{V}{t} \quad \text{--- (A)}$$

$\Rightarrow$ As per Darcy's law $Q = K i A$ --- (B)

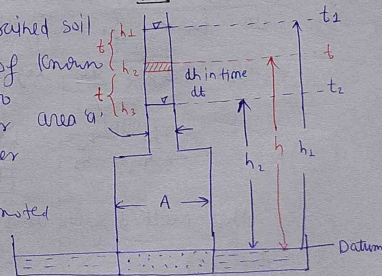
$$K i A = \frac{V}{t} \quad K = \frac{V}{t \cdot i \cdot A} = \frac{V}{t \cdot \frac{H_1 - H_2}{L} \cdot A}$$

B) Falling Head / Variable Head Permeability test

$\Rightarrow$ Suitable for fine grained soil

$\Rightarrow$ An open stand pipe of known area 'a' is inserted into the soil mass and water is allowed to flow under variable head

$\Rightarrow$ Head in stand pipe is noted at different time



→ Let in time dt head fall by ' dh ' under the head of h

→ Volume of water flows in time dt (dV) = $q \cdot dh$

→ Discharge through soil mass in time dt ,

$$q = \frac{dV}{dt} = - \frac{q \cdot dh}{dt} \quad \text{--- (A)}$$

→ As per Darcy's law $q = K i A = K \times \frac{h}{L} \cdot A$ --- (B)

from (A) and (B) $K \cdot \frac{h}{L} \cdot A = - \frac{q \cdot dh}{dt}$

$$\frac{K \cdot A}{L} \int_{t_1}^{t_2} dt = - q \int_{h_1}^{h_2} \frac{dh}{h}$$

$$\frac{K \cdot A}{L} (t_2 - t_1) = - q \ln h \Big|_{h_1}^{h_2} \quad (\because t_2 - t_1 = t \text{ time interval})$$

$$\frac{K \cdot A}{L} \cdot t = - q (\ln h_2 - \ln h_1) = q \ln \left(\frac{h_1}{h_2} \right)$$

$$K = \frac{q \cdot L}{A \cdot t} \ln \left(\frac{h_1}{h_2} \right)$$

$$K = 2.303 \frac{q \cdot L}{A \cdot t} \log_{10} \left(\frac{h_1}{h_2} \right) \quad ***$$

Note Let in given time interval t (height of water) intended pipe falls from h_1 to h_2 and in the same time q falls from $(S_1$ to $S_2)$ then.

$$K = 2.303 \frac{qL}{At} \log \left(\frac{h_1}{h_2} \right) = 2.303 \frac{qL}{At} \log \left(\frac{h_2}{h_3} \right)$$

$$\frac{h_1}{h_2} = \frac{h_2}{h_3}$$

$$h_2^2 = h_1 h_3$$

$$h_2 = \sqrt{h_1 h_3}$$

This method is used to check consistency of readings

© Capillary permeability test (Not too much important)

This test is used to find permeability of medium and capillary rise in it. It is suitable for partially saturated soil.

→ Partially saturated soil is placed in cylindrical glass tube and horizontal distance travelled by the water (x) at different time is noted.